

# Results of a new methodology for the analyze the climate change uncertainty on a high step mountain range. Application at the Andes

## A methodology with independence of model resolution and following criteria

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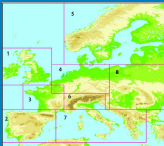
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ICRC-CORDEX, Beijing, 2019/x/14-18

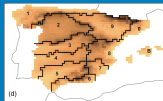


# Currently

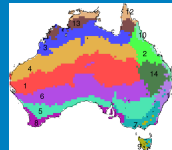
- **Space averaging:** comparison of different data-sets by spatial averaging over areas
- Definition of spatial areas different possibilities:



lon,lat boxes: artificial 'mathematical' boundaries [Christensen, 2007, *Climatic Change*]



orography: river basins [Herrera, 2010, *J. Geophys. Res.*]

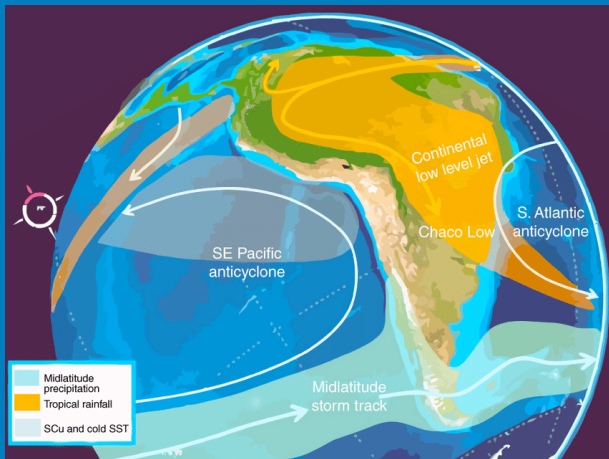


climate: climate coherence of regions [Fita, 2016, *Clim. Dyn.*]

- Following only one criteria
- **spatial incoherence** mixture of: different climates, atmospheric-dynamics, morphological characteristics...

# Circulation patterns over South America

- Climate in South America dominated by Andes mountain range, with tropical precipitation North, and storm tracks South



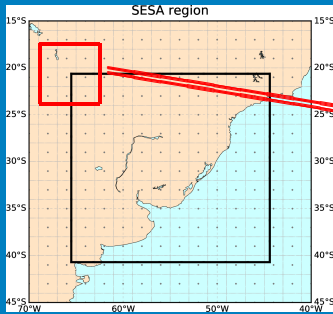
# Proposal

- Reduction of **spatial incoherence**: definition of regions using combination of different objective criteria
- Example on Andes study (3 criteria):
  - 1 **latitudinal bands**: climate coherence
  - 2 **orographic height**: climate coherence
  - 3 **mountain range face**: atmospheric-dynamics coherence
- Simultaneous application of 3 criteria to define spatial regions to spatially average model output
- Definitions:
  - criteria: definition of categories to select grid points
  - slice: union of criteria
- **spatial average**: spatial weighted average by areal overlapping of **slice** (as a polygon) and model **grid cell** (as lon,lat box)

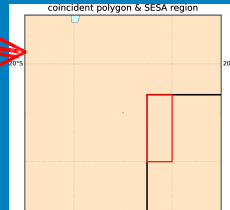


# Polygonal overlapping area

- Polygonal area overlapping: grid selection by spatial criteria.
- e.g. slice: SESA area
- $\forall$  grid cell  $((i,j))$  area  $(\mathcal{P}_g)$  within slice  $(\mathcal{P}_S)$ :



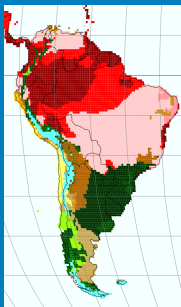
- 1 Overlapping of polygons: individual grid cells and slice. Take cell **fractions**



- 2 spatial averaging **weights** for grid cell as covered fraction of the total area of the slice  $w_g = \frac{\mathcal{A}_g^s}{\mathcal{A}_S}$

# 1st criteria: Latitudinal bands

- General climate latitudinal dependence (dist. from Equator & Poles)
- South American climate N-S bands:

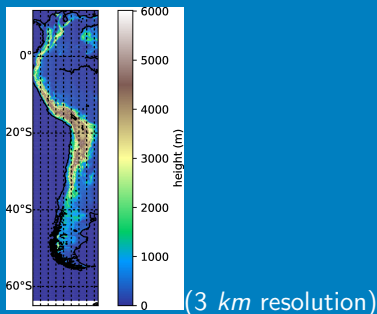


[Kottek, 2006, Meteorol. Z.]

- **criteria:** bands of latitude, from S 63.° to N 19.° every 2.° [41 categories]

## 2nd criteria: Topographical height

- Climate dependence on topographical height
- South American orography:

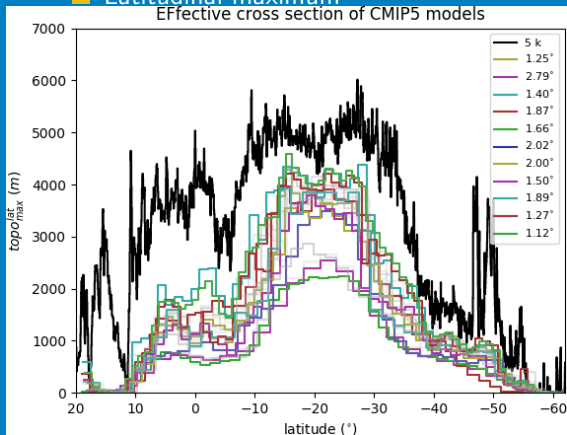


- **criteria:** height ranges of orography from 500. to 7000. m every 500. m [13 categories]

# Topographical representation CMIP5 models

■ **Challenge:** results using CMIP5 data at various poor resolutions

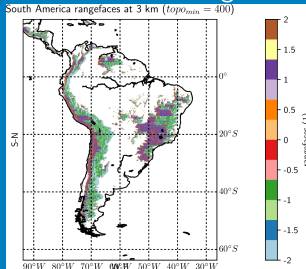
■ **Latitudinal maximum**



Loss of main peaks  
Differences about 2000 m

### 3rd criteria: mountain face

- Atmospheric dynamics dependence on wind/lee-ward side of the Andes mountain range
- Use new diagnostic **rangefaces** to define west, peak and east regions along a mountain range
- South American 'rangefaces':



(3 km resolution)

{ 2 uphill  
1 uphill valley  
0 peak  
-1 downhill valley  
-2 downhill

- Use only Andes range (single 'range' value)
- **criteria:** grid point 'uphill', 'peak' or 'downhill' [3 categories]

# rangefaces: mountain range face diagnostics

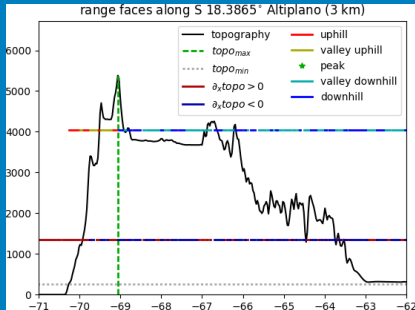
- Grid point side along longitudinal/meridional sections. e.g. Andes: latitudinal sections

- 1 *range*: grid point as part of a mountain range if it is above 400 m ( $topo_{min}$ )
- 2 *West/East*: 'x' derivative of the topography: West (**uphill**, [2]), East (**downhill**, [-2])

$$\partial_x topo(i,j) = \frac{topo(i+1,j) - topo(i,j)}{\delta x} \begin{cases} > 0. & \text{uphill (2)} \\ < 0. & \text{downhill (-2)} \end{cases}$$

- 3  $topomax(j)$ : topo. max along lat section. Defines the two major areas: uphill (before), downhill (after)
- 4 **peak** [0]: grid point  $imax$  where  $topomax(j)$
- 5 *valleys*: local minimums: **uphill valleys** [1] and **downhill valleys** [-1]
- 6 Repeated along all meridional sections per mountain range (contiguous grid points above  $topo_{min}$ )
- 7 Unique identification per mountain range

# rangefaces: mountain range face diagnostics

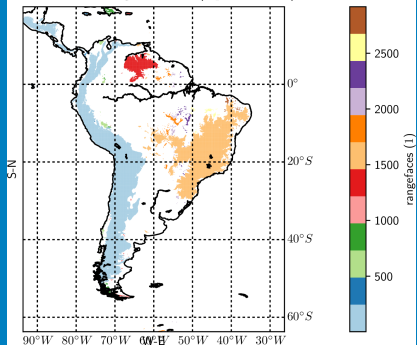


(e.g. section lat. S 18.3865°, 3 km res.)

rangefaces

- 2 uphill
- 1 uphill valley
- 0 peak
- 1 downhill valley
- 2 downhill

South America ranges at 3 km ( $topo_{min} = 400$ )

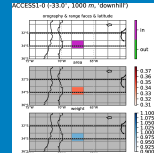
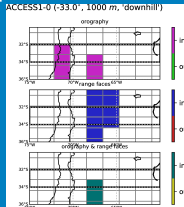
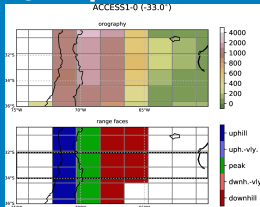


Result on:

- 3km resolution topography
- y-axis section
- $topo_{min} = 400\text{ m}$

# Slicing: regions following different criteria

- e.g. Andes mountain range following: lat. bands, orog. and range-face. Total pot. slices:  $41 \times 13 \times 3 = 1599$
- e.g. CMIP5 ACCESS1.0 for: latitudinal band= $[S\ 34^\circ, S\ 32^\circ]$ , height= $[750. m, 1250. m]$  and 'downhill' range face:  $[-2.5, -0.5]$



- Final area ( $\mathcal{A}_{C1,C2,C3}$ ); acc. product of area of each criteria
- weights ( $\mathcal{W}$ ): fraction of slice area ( $\mathcal{A}_{C1,C2,C3}$ ) covered by each grid point  $\mathcal{A}_{C1,C2,C3}^g[k]$

$$\mathcal{W}[k] = \frac{\mathcal{A}_{C1,C2,C3}^g[k]}{\mathcal{A}_{C1,C2,C3}}$$

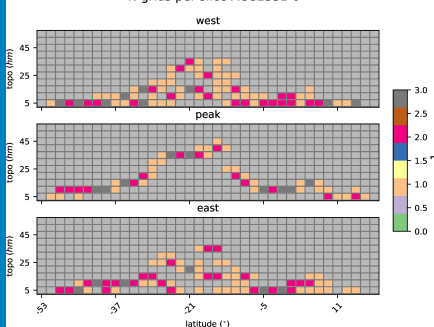
$$\mathcal{A}_{C1,C2,C3} = \sum_{k=1}^{N_{1,2,3}} \mathcal{A}_{C1,C2,C3}^g[k]$$



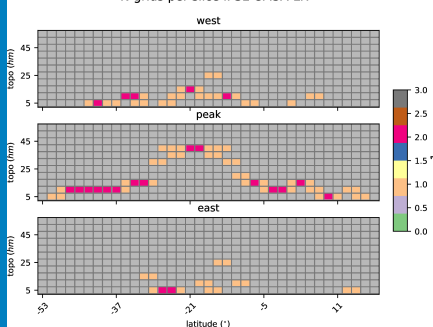
# Slicing: description

## ■ Number of grid cells per slices: ACCESS 1.0 & IPSL-CM5A-LR

N grids per slice ACCESS1-0



N grids per slice IPSL-CM5A-LR



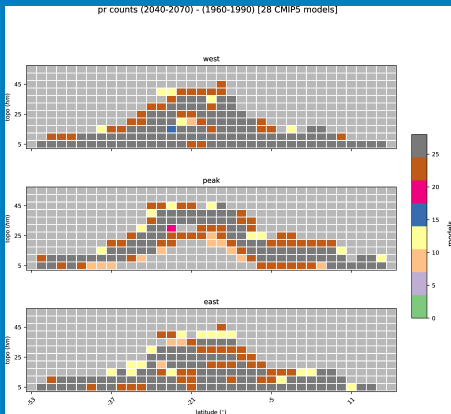
Number of grid cells per slice depend on model resolution:

- ACCESS 1.0:  $1.875 \times 1.25^\circ$
- IPSL-CM5A-LR:  $3.75 \times 1.895^\circ$

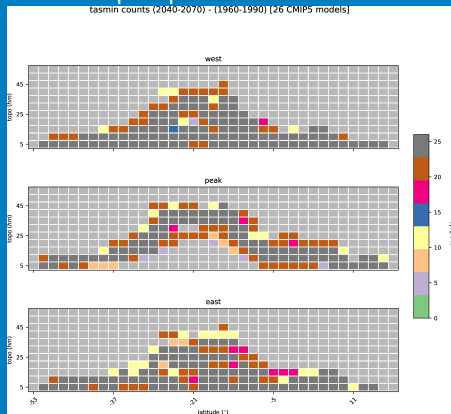
# Slicing: description (cont. ii)

## ■ Number of models with data per slice for precipitation and tasmin

pr counts (2040-2070) - (1960-1990) [28 CMIP5 models]



tasmin counts (2040-2070) - (1960-1990) [26 CMIP5 models]



Each slice with different models and ensemble size

# Uncertainty of climate change signal

- For each slice, response to climate change (*fut*, 2040-2070 vs *hist*, 1960-1990) for: pr, tasmin, tasmax is computed. ESGF monthly data (28 models pr, 26 models tasmin/tasmax)
- Uncertainty of response is computed as in [Kendon et al., 2008, J. Climate] (after [Storch and Zwiers, 1999, ] )

1 signal-to-noise-ratio:  $SNR = \frac{\langle ENS_{resp} \rangle}{\left( \frac{\sqrt{(\sigma_{yrvar}^{hist})^2 + (\sigma_{yrvar}^{fut})^2}}{2} \right)}$

- 2 significance at 5 % by two-tailed T-student:

1 equivalent T-student:  $t_{equiv} = SNR \sqrt{\frac{N_{mod}}{2}}$

2  $\gamma$ -degrees of freedom:  $\gamma = (N_{mod} - 1) \left[ \frac{\left( 1 + \left( \frac{\sigma_{yrvar}^{fut}}{\sigma_{yrvar}^{hist}} \right)^2 \right)^2}{1 + \left( \frac{\sigma_{yrvar}^{fut}}{\sigma_{yrvar}^{hist}} \right)^4} \right]$

3 finally, significance:  $\begin{cases} t < t_{theo}(\gamma, 0.025) \\ t > t_{theo}(\gamma, 0.975) \end{cases}$

## Uncertainty of climate change signal (cont.)

- SNR and significance test computed from variances ( $\sigma_{yrvar}^{hist/fut}$ ):

- interannual variability of the ENS:

$$\frac{1}{N_{mod}} \sum_{imod=1}^{N_{mod}} (yrvar_{imod}^{hist/fut})$$

- ENS spread of the interannual variability:

$$\sigma(yrvar_{imod}^{hist/fut})$$

- ENS spread of the annual mean:

$$\sigma(yranmean_{imod}^{hist/fut})$$

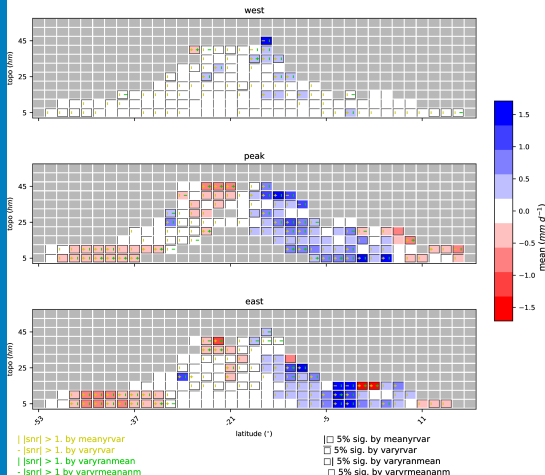
- ENS spread of the interannual anomalies:

$$\sigma(yrmeananom_{imod}^{hist/fut})$$

# Results

## ■ ENSmean precipitation response

pr response (2040-2070) - (1960-1990) [28 CMIP5 models]



- Precipitation increase mostly East-tropical regions
- pr decrease everywhere else
- certainty and significance slice and variance dependency

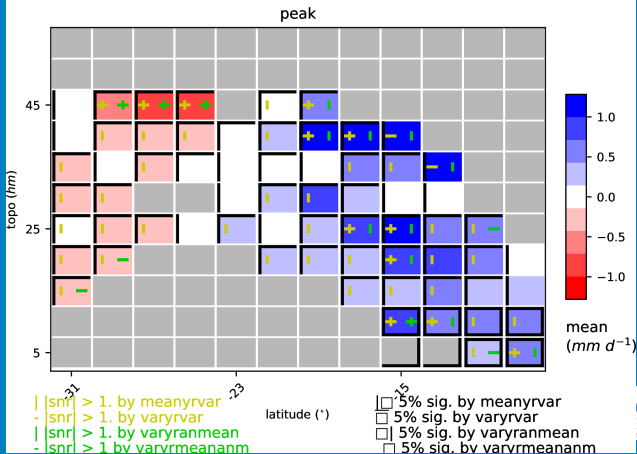
'+' : certain by  $|SNR| > 1$

'□' : T-student sig.

# Results

- Zoomed example on slicing for pr: 'peak'

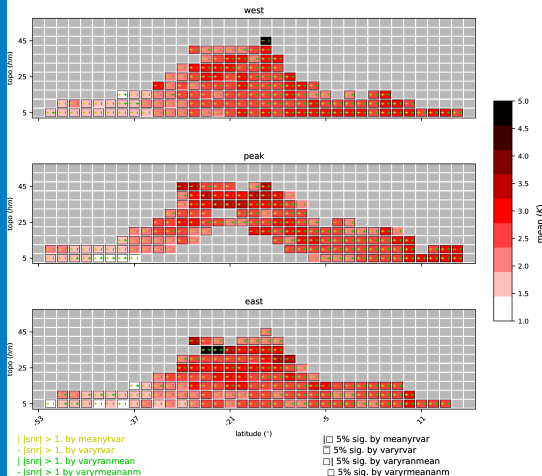
ZOOM pr response (2040-2070) - (1960-1990) [28 CMIP5 models]



# Results

## ■ ENSmean minimum temperature response

tasmin response (2040-2070) - (1960-1990) [26 CMIP5 models]

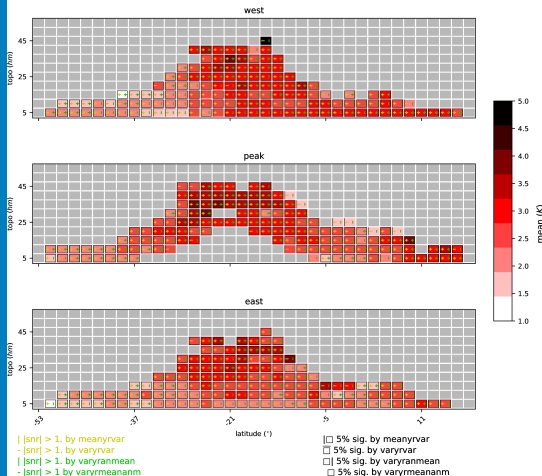


- stronger increase at higher altitudes, 'peak' and 'east' areas
- no signal at high latitudes
- certainty and 5% significance slice and variance dependency

# Results

## ■ ENSmean maximum temperature response

tasmax response (2040-2070) - (1960-1990) [26 CMIP5 models]



- Stronger increase that tasmin
- Increase also at high latitudes
- Slight decrease in certainty and significance



# Conclusions

- Analysis of climate change signal of CMIP5 by **slices** (latitude, orography and mountain side) over the Andes
- Objective definition of **coherent** spatial areas **no-interpolation**
- latitudinal, height and mountain side **dependency** of climate signal on **precipitation**
- **Wider thermal** response, stronger at higher altitudes
- Stronger response of **tasmax**
- Strong **dependency** of signal certainty and significance on **slice** and method estimation of uncertainty

Thank you for your attention !!

# The Andes - South America's backbone for interdisciplinary studies

- Andes comprise a heterogeneous territory in which *climate variations affect socio-economic activities and could endanger vulnerable communities*
- The Andes are a biodiversity hotspot hosting 10% of the world's plant species, and providing *critical ecosystem services* as water provision, soil protection, carbon storage
- Vulnerability of the Andean population is high and therefore *increasing knowledge of climate change is crucial to guide adaptation policies*



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**ANDES** are a gap! and an **opportunity to CORDEX**

→ linking with **Andex community** under the umbrella of **GEWEX**

Thank you!

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